ISOTOPIC AND ELEMENTAL COMPOSITION OF CHROMIUM, IRON, AND NICKEL IN TYPE I DEEP-SEA SPHERES: S. Xue^{1,2}, C.-Y. Shih³, H. Wiesmann³, D.E. Brownlee⁴, G.S. Hall¹, G.F. Herzog¹, and L.E. Nyquist⁵; ¹Dept. Chemistry, Rutgers Univ., New Brunswick NJ 08903, USA, ²Grad. School Oceanog., Narragansett Bay Campus, Univ. Rhode Island, Narragansett RI 02882-1197, USA, ³Mail Code C23, Lockheed-Martin, 2400 NASA Road 1, Houston TX 77258, USA, ⁴Dept. Astronomy, Univ. Washington, Seattle WA 98105, USA, ⁵NASA Johnson Space Center, Houston TX 77058, USA.

Summary - We report analyses of Cr, Fe, and Ni in type I deep-sea spheres with masses ranging from 43 to 246 µg. The isotopic compositions of Cr and of Fe were measured by TIMS; the elemental concentrations of Cr, Fe, and Ni and the isotopic compositions of Ni were measured by ICP-MS. In most spheres evaporation has led to significant and similar average degrees of mass-dependent fractionation, Φ , for all three elements. The average value, $\langle \Phi \rangle \sim 15$ %/AMU, corresponds to mass losses of 80-85% assuming open-system evaporation of the atoms. We find that 1) $\Phi_{N_i} \propto (\Phi_{Fe})^n$, $n \sim 3$, perhaps because Ni evaporates more slowly than Fe but faster than FeO; and 2) $\Phi_{Cr} \sim \Phi_{Fe}$ for all Φ_{Fe} .

Background - Type I deep-sea spheres consist mainly of O, Fe, and Ni, which may occur as magnetite, wüstite, hematite, and/or metal, and of other elements in smaller concentrations. Kyte [1] pointed out that Cr/Fe ratios are diagnostic for certain kinds of meteoritic metal and therefore could help in identifying parent materials for the type I spheres. Of late it has become clear that heating of type I spheres during atmospheric entry caused large evaporative losses and, along with them, mass-dependent fractionation of the isotopic abundances of O, Fe, and Ni [2-6]. To the extent that such losses differ for each element, they must be known before the composition of a sphere as received can be related to that of the parent material. Similarly, the losses must be known in order to relate the observed masses to the preatmospheric masses. Under certain conditions the extent of isotopic fractionation of an element can furnish a measure of the degree of loss. In a study of eight type I spheres Davis and Brownlee [4] found that Φ_{Fe} does not equal Φ_{Ni} . Nyquist et al. [7], in oral presentation, reported isotopic data for Cr, Fe, and/or Ni in four spheres. We expand on this earlier work in an effort to learn more about the initial sizes and elemental ratios of type I spheres.

Experimental methods and results - We used the methods of [7] to measure the concentrations and isotopic

Table 1. Average compositions

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Sample	Cr	Fe	Ni
		(‰)	
This work	1.6	725	35
Ref. [8]	1.2	688	14.2
Ref. [1]	1.2	649	41

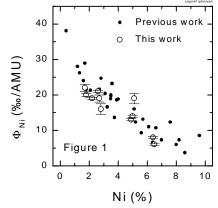
ratios of Cr, Fe, and Ni in type I spheres. From the isotopic ratios we obtained Φ by calculating δ -values, plotting them vs. mass, and reading the slope. Table 1 compares averages of the new elemental concentrations and those of [7] ("This work") with literature values. Most of the Fe contents fall in the range between 60-75% by mass; Cr and Ni contents vary by well over a factor of ten just as in [1] and [8]. Among our 13 spheres, 2 contained very little Ni (<0.1%); in ref. [8], 16/57 spheres have very low Ni contents. Kyte [1, Table 1], reported Ni contents below a

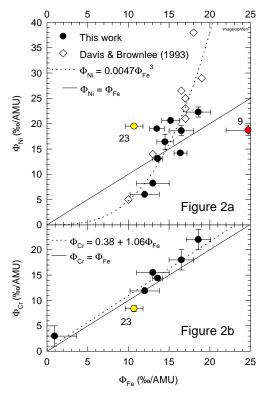
detection limit of ~0.5 for 2 of 15 spheres. In light of [9], we hesitate to write off these low-Ni objects as terrestrial but treat them separately. As for Cr, the range 0.16 to 4.03 % includes all

samples except the outlier KK1-23 (16.4% Cr).

In agreement with previous work for whole spheres [6], we find that Φ_{N_i} correlates inversely with Ni (Figure 1). Similar but weaker trends can be seen for $\Phi_{\rm Fe}$ and $\Phi_{\rm Cr}$ vs. Ni content if one outlier (KK1-9) and one sphere with unusually high Cr (KK1-23) are excluded.

Although the average degree of mass-dependent fractionation (%/AMU) is similar for Cr (15.0±4.7), Fe (15.3±5.6), and Ni (16.1±5.2), the detailed relations differ. Figure 2a confirms





the observation of [4] that Φ_{Ni} at first increases more slowly and then more rapidly than Φ_{Fe} . Figure 2b shows that Φ_{Cr} increases at about the same rate as Φ_{Fe} ; the data are consistent with a one-to-one linear relationship.

Discussion - The trends in Figure 2 can be understood if the evaporation rate of Fe initially exceeds that of Ni and then slows as metallic iron forms a less volatile oxide. oxygen isotopic data suggest that evaporation of iron oxide remains appreciable, however. Average values of oxygen fractionation, Φ_0 , measured for comparable or greater numbers of spheres of various sizes [2] and recalculated relative to air assuming no kinetic isotope effect during oxidation of iron, are as follows: $\Phi_{O-air} = 12.4$ (~400µm); 9.6 (240-340 μ m); and 8.4 (100-150 μ m). If we model the evaporation of Cr, Fe, and Ni as an open system distillation of the monoatoms and of oxygen as a species with a fractionation factor for ¹⁷O/¹⁶O of 1.0110 [10], then the average values of Φ , $\langle \Phi \rangle$, correspond to losses of 80-85% for the transition metals but only about 60% for oxygen ($\langle \Phi_{O} \rangle = 9.6$). Use of the Rayleigh equation shows that the accompanying loss of 60% of the oxidized iron could

account for ~8 of the 15 ‰/AMU of $<\Phi_{Fe}>$. The other 7 ‰/AMU of $<\Phi_{Fe}>$ may result from early and more rapid evaporation of Fe as metal. We note that the fit of Figure 2a turns upward $(d\Phi_{Ni} \ / \ d\Phi_{Fe} \ge 1)$ at $\Phi_{Fe} \sim 8$ ‰/AMU.

The similar values of Φ_{Cr} and Φ_{Fe} imply that measured Cr/Fe ratios, which average ~0.002, changed little during sphere formation; evidently Cr behaved chemically much as did Fe. Figure 3 compares Cr/Fe ratios for type I spheres (bars) [1,7,8] with average bulk values for selected meteorite groups. (Meteorite group names are plotted at the appropriate Cr/Fe ratios, but arbitrary heights). If Type I spheres form from chondritic matter, then the process fractionates Fe and Cr. The Cr/Fe ratios in kamacite from C-chondrites, ~0.001-0.006 [11] (not plotted), are a fair match to the observations.

The isotopic data indicate average evaporative losses of 80-85% for Cr, Fe, and Ni. Observed type I sphere masses must therefore be increased by at least a factor of 5 to obtain initial masses, and by

more if ablative losses occurred (see also [12]). The decrease in diameter due to evaporation is counterbalanced by the formation of lower density oxide from higher density metal.

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